



GEOCHEMICAL EVALUTION FOR FORMATION WATER OF SOME WELLS IN AMAL OIL FIELD, SIRT BASIN, LIBYA

Asmaa D. Alnajar¹, Naima. K. Elgariani², Anwar M. Alsaiah³

^{1,2,3}University of Tripoli, Geological Engineering Department, Libya

Alnajarasmaa733@gmail.com , nkgariani@gmail.com

Abstract

The present paper deals with classification of oilfield water of Maraghand Amal formations at Amal oil field, concession NC-12 ,in the eastern margin of the Sirt basin, Libya.Ten chemically analyzed water samples from different wells were used in this study; chemical analyses of the major ions of the water samples showed Total Dissolved Solids (TDS) in the ranges from 205487 to 234912mg/l.

The geochemical ratios of Na/Cl, ((Na-Cl)/SO₄), and ((Cl-Na)/Mg) were calculated for formation water of ten wells classification following Sulin's (1946), whereindicate that all of these water are Cl-Ca type waters under subsurface conditions; this type reflects isolated waters with hydrostatic conditions.According to Bojarski's (1970), the chloride-calcium type of this water subdivided to chloride-calcium III, which reflects the favorable zone for the preservation of hydrocarbons, followed by chloride-calcium II in other wells.This class reflects the transition zone between active hydrodynamic and stable zones.

Graphical methods in water analysis enable efficient pattern comparison for quick identification of differences between water samples.Consequently, both the piper diagram and the stiff diagram were utilized in this study. The piper diagram in this study indicates most of the samples were classified as (Na⁺+K⁺) and Ca²⁺ types of cations and Cl⁻ types of anions;then these waters consist of the alkalis that exceed alkaline earths.The Stiff Diagram polygons showed most of the samples reveals similar chemical compositions with high chlorine and sodium levels, followed by magnesium and bicarbonate, while sulfate is low. In contrast, the waters in some wells have higher magnesium concentrations.

Keywords: Geochemistry; Classification; Amal Field; Oil field Waters.

المستخلص

يتناول هذا البحث تصنيف مياه حقول النفط في تكويني مراغوأمال في حقل مال النفطي، امتياز NC-12، فيا لهامش الشرقي لحوض سرت، ليبيا. تم استخدام عشر عينات من المياه المحللة كيميائياً من آبار مختلفة في هذه الدراسة؛ أظهرت التحاليل الكيميائية للأيونات الرئيسية في عينات المياه أن المواد الصلبة الذائبة الكلية (TDS) تتراوح بين 205487 إلى 234912 ملغ/لتر.

تم حساب النسب الجيوكيميائية Na/Cl و $(Na-Cl)/SO_4$ و $(Cl-Na)/Mg$ لمياه التكوين لعشر آبار وفقاً لتصنيف سولين (1946)، حيث تشير إلى أن جميع هذه المياه من نوع $Cl-Ca$ تحتظر وفتحت السطحية؛ هذا النوع يعكس مياهًا معزولة تحتظر وفهيدروستاتيكية. ووفقاً البوجارسكي (1970)، تم تقسيم الماء من نوع كلوريد-كاليسيوم إلى كلوريد-كاليسيوم III، الذي يعكس منطقة مواتية لحفظ الهيدروكربونات، يلي هكلوريد-كاليسيوم II في آبار أخرى. تعكس هذه الفئة منطقة انتقالية بين المناطق هيدروديناميكية نشطة ومناطق مستقرة.

تمكن الطرق البيانية في تحليل المياه من مقارنة الأنماط بكفاءة لتحديد الفرق بين عينات المياه بسرعة. وبالتالي، تم استخدام كل من مخطط بايبر ومخطط ستيف في هذه الدراسة. يوضح مخطط بايبر في هذه الدراسة أن معظم العينات تم تصنيفها علناً بأنها من عالكاتيونات $(Na^+ + K^+)$ و Ca^{+2} والأيونات من نوع Cl ؛ وبالتالي فإن هذه المياه تتكون من القلويات التي تفوق القلويات الأرضية. أظهرت مضلعات مخطط ستيف أن معظم العينات تكشف عن تراكيب كيميائية متشابهة بمستويات عالية من الكلور والصوديوم، تليها المغنيسيوم والبيكربونات، بينما يكون مستوى الكبريتات منخفضاً. على النقيض، تحتوي المياه في بعض الآبار على تراكيز أعلى من المغنيسيوم.

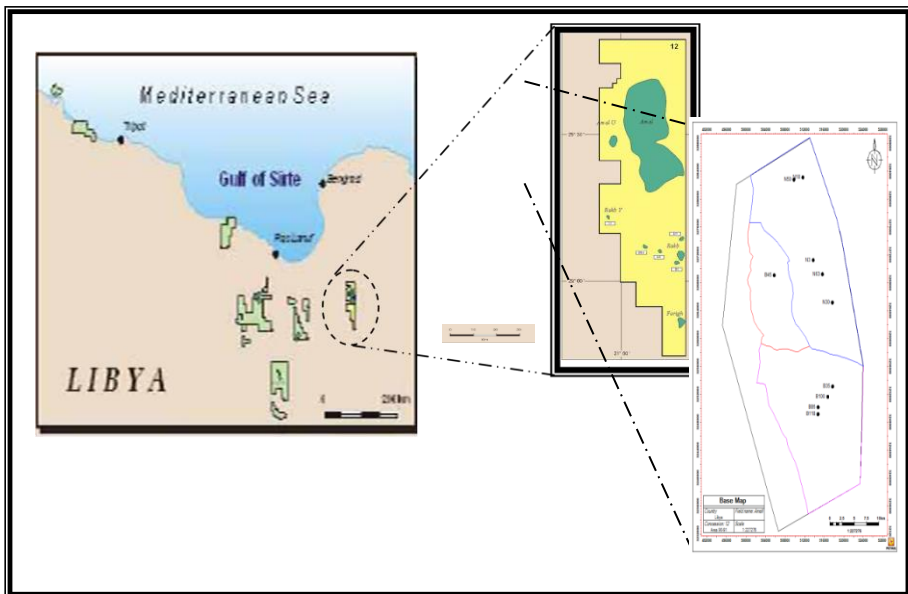
1. Introduction

Water geochemistry provides a series of powerful tools for solving various oilfield development and production problems. Specifically, naturally occurring chemical "tracers" in water can be used to identify the origin and track the movement of water in oil fields, as well as to predict the precipitation of scales. These naturally occurring tracers include the absolute and relative abundance of the various dissolved salts (anions and cations) as well as the isotopic composition of certain cations and the hydrogen and oxygen stable isotopic composition of the water itself. Which formation is responsible for water moving behind casing can be determined by comparing formation water geochemistry data with data from produced waters.

The produced water or petroleum-associated waters often have salinities in the range from less than 10,000 up to more than 200,000 mg/l. However, most oilfield brines have salinities much higher than seawater salinities. Hence, oilfield waters are often termed oilfield brines, especially when their total salinities are more than 100,00 mg/l (Refai, 2011).

1.1 Location of Study Area:

The Amal Field is situated in Concession 12, on the eastern margin of the Sirt Basin, in Cyrenacia Province, Libya. The concession lies some 275km southeast of the Mediterranean coast and 925km ESE of Tripoli. Geographically, Concession 12 lies in desert terrain, hilly in the north and sandy in the south. The nearest town is Augila, which lies immediately to the south of the field. The Amal Field is located on Amal High and Maragh Low. It lies in the northeast flank of Sirte Basin in the northern part of Concession 12, between coordinates of Latitudes 29°19' 04" and 29° 41' 47" North and Longitudes 21° 01' 05" and 21° 14' 35" East, Figure (1). (Haruge company, 2005)



**Figure (1) Illustrate Location map of the concession 12, Libya
modified after Haruge company.**

1.2 Literature review

The name Amal Formation has recently been used for a formation in the subsurface of the Amal Field area (Roberts, 1970). Therefore, the Amal Formation is here proposed as a new formation in the subsurface of the eastern Sirt Basin, however, it is identical to the rock unit used by Roberts. (1970) provides a valuable discussion of the lithology of this formation in the Amal Field, its regional structural setting, and

its importance as a major petroleum reservoir. The Amal Field was discovered by well B1-12, drilled in 1959, when Mobil first drilled the Amal Field.

2. Stratigraphy and petroleum system

2.1 Geological and Structural Setting

The Sirt Basin is a large sedimentary basin located in central north Libya that has produced significant oil and gas resources over several decades (Schlumberger service company, 1997). Formed during the Late Paleozoic to Mesozoic breakup of the supercontinent Pangaea, the basin transitioned from an extensional rift setting to a passive margin as the African plate migrated northwards relative to Eurasia (Jiranek, et al., 1998). The Sirt Basin is a major intracratonic rift system in central Libya. The sedimentary succession within the basin reflects its tectonic and structural evolution, which is closely related to the opening of the Atlantic Ocean and the opening and subsequent closure of the Tethys Oceans in Mesozoic and Tertiary times (Whitbread, 1989). In general, the sedimentary sequence of the Sirt Basin overlies Precambrian igneous and metamorphic basement rocks and comprises sporadically distributed and poorly dated Lower Paleozoic clastics which are unconformably overlain by a thick Lower Cretaceous and Tertiary clastic and carbonate sequence.

2.2 Lithology of study area:

The stratigraphic nomenclature used is that employed by Veba Oil Operations (previously Mobil Oil Libya). The stratigraphic column in the study area comprises deposits that range from pre Cambrian to Tertiary presented in Figure (2).

2.2.1 Basement:

The basement has been described by Mobil (1979) as consisting of two major igneous centres. A widespread Cambro-Ordovician or older (500 million years) basement, consisting predominantly of granitic rocks with rhyolites and tuffs as well as trachytic volcanics. This is the type of basement penetrated by the wells in the southern part of the field. To the north, Lower Cretaceous-Jurassic (135 million years) intermediate volcanics (F1-12) are centred on the Tharwa High (North Amal).

2.2.2 Amal Formation: Cambro-Ordovician

The Amal Formation is predominantly a sandstone sequence. It is heterogeneous in color, grain size and sorting. The grain size ranges

from very fine sand to cobbles with the medium to coarser grain sizes being more common. Accessory constituents are feldspar, mica, pyrite, hematite and various dark minerals. Clays, sericite and rarely dolomite are found as cementing materials in much of the formation. Interbedded with the sandstone, but comprising a much lesser part of the formation, are gray silty clays and gray, green and red, brittle, micaceous shales. In addition to the sedimentary rocks, volcanic rocks in the form of dikes, sills or flows are found at a number of horizons in the upper part of the formation. Contact relationships in the B1-12 well, the Amal Formation is unconformably overlain by the Maragh Formation. The base of the formation was not reached in the B1-12, but in other wells, such as B2-12, the Amal Formation unconformably overlies volcanic rocks. In the type section, the top of the Amal Formation is placed at the change from the friable dolomitic sandstone of the Maragh Formation to the more firmly cemented quartzitic Amal sandstone. Because of onlapping relationships around the Amal high, the Amal Formation can be overlain by various Rakb Group clastics or carbonates. The base of the formation, where reached, occurs at the change from sandstone to volcanic rocks. Paleontology and Age. Radiometric determinations on volcanics within and beneath the Amal Formation have given a broad and inconsistent spread of ages ranging from Cambrian to Jurassic. Regional considerations, however, suggest that the most likely age of the Amal Formation is Cambro-Ordovician. Origin of Name This formation is the principal reservoir of the Amal Field from which it derives its name, and its importance as a major petroleum reservoir. Distribution the Amal Formation, or its equivalents, are widely spread in the eastern Sirt Basin. it may be equivalent to the Hofra Formation in the western Sirt Basin. In the Amal Field area, the Amal Formation varies greatly in thickness.

2.2.3 Maragh Formation: Upper Cretaceous

The Maragh Formation lies unconformably on the Amal Formation or the Barremian Sand/ Nubian Sand, separated from the underlying formations by the Sirt Unconformity, the Maragh Formation is referred to as the Bahi Formation.

The Maragh Formation is thinning from northeast to southwest. The presence of the Maragh in this position can be explained as either transgression up the lows or a preferential preservation in the low protected areas with erosion in the higher areas. The distribution of the Maragh Formation suggests that it was once more widely distributed

and that what is present now is an erosional remnant, erosion could have happened at both the Santonian and Campanian unconformities. The Maragh Formation is easily distinguished from the underlying Amal Group by its relatively soft and unconsolidated state.

The Maragh Formation was divided into three units by Geopec (1994), and this division has been followed:

- The Maragh unit X, which is represented by sequences of loosely consolidated conglomerates.
- The Maragh unit A, followed by less conglomeratic than Maragh X, is described as being friable, very fine to fine grained moderately to poorly sorted, glauconitic sublitharenite to lithic subarkose, accompanied by phosphatic and bioclastic remains, interbedded with mudstone/shale and carbonate layers.
- The Maragh unit B is distinguished by the calcite and dolomite cemented sandstones underlying less cemented sandstones of the Maragh A and the overlying shales and limestones of the Rakk C.

The Amal Field covers an area of about 880km² and is situated on the Amal High, on the eastern flank of the Sirt Basin, adjacent with the Rakk and Nafoora-Augila fields. The Amal Field is flanked by the Ajdabiya Trough to the west and the Maragh Low to the east. The structure plunges to the north, into the Tharwa Low and the reservoir pinches out, up-dip, to the south. The crest of the structure lies near the southern boundary of the field at a depth of about -9,530' TVDSS. With a nominal oil water contact of -10,260' TVDSS. The Amal Field has been subdivided into three sectors: B, N and E. Both B and E sectors are on the footwall to the major intra-field fault which separates it from the N sector in the hanging wall. B and E sectors are separated from one another by a high-angle east-west oriented fault. Figure 3.

	Period	Sub Period	Group	Formation	Member	Lithology	Note			
Tertiary		Post Miocene	Najah	Garet Uedda			Loose surface sands			
		Miocene		Giarabub			Dolomites and limestones interbedded with clays			
		Oligocene		Gebel Akhdar	Arida		Sandstone and limestone interbedded with mudstone			
		Eocene	Cyrenaica	Gialo	Augila		Limestone interbedded with sandy clay			
					Rashda		Limestone with minor inter-bedded limestone and sand			
					Smara		Argillaceous limestone with clay			
		Buddaffar				Argillaceous limestone with clay				
		Paleocene	Farigh		Etta		Limestone with dolomite			
					Mesdar		Limestone becoming dolomite downwards with dolomite at base			
					Abu Fas		Argillaceous marls and calcareous shales with thin interbedded limestones			
Upper Sabil Sheterat					Limestone, shaly with hard dolomite					
Mesozoic		Upper Cretaceous	Harash	Rakb	Rakb A		Shale interbedded argillaceous limestones and siltstone			
					Rakb B		Shale and limestone with minor interbedded sandstones and siltstones			
					Rakb C		Shale with interbeds of limestone and dolomite. The base of the unit is often formed of silty sandstone			
					Rakb D		Shale grey-black, interbedded sandstone, white to very fine grained			
					Maragh B		Caliche and dolomite cemented sandstone			
					Maragh A		Obscure sandstone, fine grained interbedded with mudstone/shale			
					Maragh X		Loosely consolidated conglomerates			
					Unit 1		Multicolored shale with interbedded silty quartzite sandstone			
					Unit 2		Argillaceous quartzite sandstone with interbedded red shale and siltstone			
		Lower Cretaceous			Nubian	Unit 3		White quartz sandstone, fine to medium grained		
Unit 4						Interbedded multi-colored quartzite sandstone, siltstone and shale				
Unit 5						White to grey sandstone interbedded siltstone and shale				
Paleozoic		Upper Jurassic	Sirte	Amal	Unit V		A fine to very coarse grained conglomerate, calcareous quartzite sandstone, interbedded are calcarenites			
					Unit IV		A very fine to very coarse grained, poor to medium sorted, grey to red			
					Unit III		Multicolored quartzite sandstone, fine grained to coarse grained			
		Carboniferous					Unit II		Coarse to fine grained sandstone interbedded with thin conglomeratic	
							Unit I		Very coarse, grey to pink, light to dark quartzite sandstone	
							Unit I		Conglomerate sandstone grey to brown mobile pebbles, high gamma clay	
		Cambrian						Unit I		Greenish rocks with micaceous and felds as well as brachiopod corals
								Unit I		
								Unit I		
		Pre Cambrian							Unit I	
Unit I										

Source: [Bosch, 1927](#); The [Geological, Geophysical, and Structural History of the Margh Line and Surrounding Area](#), Eastern Margin, S. El-Saifi.

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Figure (2) Stratigraphic Columnar Section of Study Area(Haruge company, 2005).

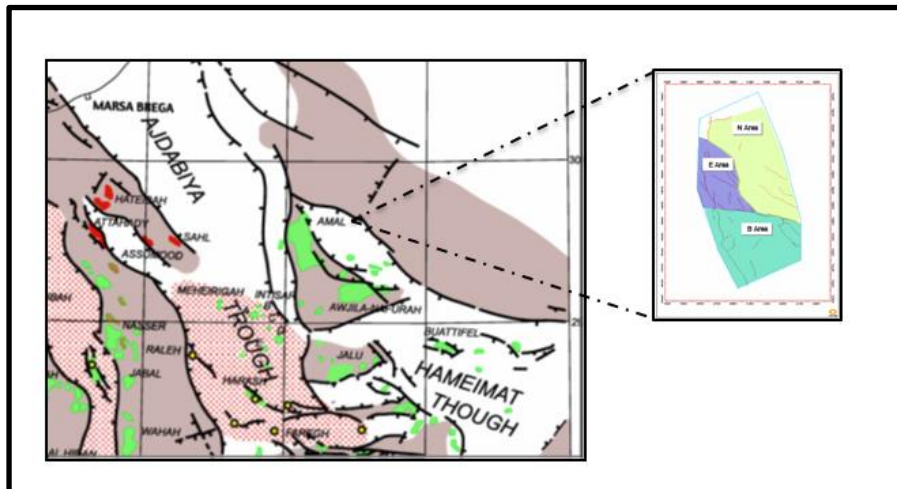


Figure (3) Structural Location of the Amal Field on The Amal High

3. Previous Studies:

Research on the chemical properties of formation water has concentrated on two primary aspects: the classification of formation water and its distribution characteristics. Many previous studies have explored the classification of formation water.

Liu et al. (2024) studied the geochemical characteristics and origin of the formation water of the saline lake basin: a case study of the Quaternary Qigequan formation in the sanhu depression, qaidam basin. Their findings revealed that the formation water in the study area has a high total dissolved solids (TDS) content and is mainly type IV and V of CaCl_2 .

Abou El Leil et al. (2022) studied classification and geological environment identification of produced water in oilfields. Moreover, they found the type of formation water is belonging to CaCl_2 type and deep marine environment.

Studied prediction of calcium carbonate and calcium sulfate scale deposits in Mabruk oilfield, Libya: studied by (Alnajjar and Refai 2021). Naima Elgariani (2007) studied the oilfield water of the Nasser oil field; according to Sulin's, Bojarski's, and Scholler's geochemical investigation, the petroleum-associated water belongs to Cl-Ca, followed by SO_4 -Na, Cl-Mg, and HCO_3 -Na waters.

4. Purpose of study:

Knowledge of oilfield water type and classification, from the chemical composition of oilfield waters in Amal oilfield, Sirt basin, Libya. In other words, determine the environment of deposition.

5. Methods of study:

The chemical analysis of ten oilfield water samples collected from: B – 35, B – 45, B – 86, B – 100, N – 3, N – 18, N – 30, N – 50, and N – 63 wells. These samples were taken from December 2018 to November 2023 and tested in the laboratory of the Amal labin Harouge Oil Operations (Harouge, 2005), which was distributed on the Amal oilfield in the Sirt basin, Libya. Figure (1). Table (2) summarizes the results of the analysis, which were calculated first in mg/l. This interpretation includes determination of total dissolved solids (TDS), major cations (Na^+ , K^+ , Ca^{2+} , and Mg^{2+}) and anions (SO_4^{2-} , Cl^- , and HCO_3^-), type of water, as well as equivalent per million (epm) Table (3), so it can be used later on in the calculation of the percentages of the different parameters in the water type classification approach. The water type classification presented here adheres to Sulin's classification

established in 1946, with modifications introduced by Bojariski in 1970, as presented in Collins (1975). This classification was selected due to its relevance and alignment with the data under analysis.

Furthermore, the Rock Work works soft were used in ternary, and finally results of these parameters were put statistically onto specified tables.

6. Geochemical Interpretations:

Geochemical interpretation was conducted using the following Formulas:

Part per million (ppm):

A dimensionless concentration term that expresses the number of unit weights of solute per million unit weights of solution (ppm).

If the water is very fresh, the specific gravity is essentially 1.0, and ppm and mg/l are equal. However, as the TDS of the water increases, the specific gravity increases, and the units become increasingly different.

$$ppm = \frac{mg/L}{S.G} (1)$$

Where:

S.G: Specific gravity of water (relative density).

Milliequivalents per Liter (meq/L):

$$meq/L = \frac{(mg/L) \times valance}{Atomic\ or\ molecular\ wieght} (2)$$

Equivalents per Million (epm):

This unit of concentration is calculated as follows:

$$epm = \frac{meq/L}{S.G} (3)$$

To Classify of oilfield waters (types of waters) according to Sulin, we used the following formulas:

In this classification, the epm is the usable unit. It's based on:

1. Na/Cl.
2. (Na – Cl)/SO₄.
3. (Cl – Na)/Mg.

I- Type of Waters is Cl – Ca:

When:

$$\text{Na/Cl} < 1.0, (\text{Na} - \text{Cl})/\text{SO}_4 < 1.0 \text{ and } (\text{Cl} - \text{Na})/\text{Mg} > 1.0. \quad (4)$$

II- Type of Water is Cl – Mg:

When:

$$\text{Na/Cl} < 1.0, (\text{Na} - \text{Cl})/\text{SO}_4 < 0.0 \text{ and } (\text{Cl} - \text{Na})/\text{Mg} < 1.0. \quad (5)$$

III- Type of Water is HCO₃ – Na:

When:

$$\text{Na/Cl} > 1.0, (\text{Na} - \text{Cl})/\text{SO}_4 > 1.0 \text{ and } (\text{Cl} - \text{Na})/\text{Mg} < 0.0. \quad (6)$$

IV- Type of Water is SO₄ – Na:

When:

$$\text{Na/Cl} > 1.0, (\text{Na} - \text{Cl})/\text{SO}_4 < 1.0 \text{ and } (\text{Cl} - \text{Na})/\text{Mg} < 0.0. \quad (7)$$

Bojarski noted considerable variation in the composition of Cl – Ca type waters and categorized these waters accordingly. Based on Bojarski's research, we utilized the following formulas:

Also, in this classification, the epm is the usable unit. It's based on the Na/Cl ratio.

a) Cl – Ca type I water:

When:

$$\text{Na/Cl} > 0.85 \quad (8)$$

b) Cl – Ca type II water:

When:

$$\text{Na/Cl} = 0.85 - 0.75 \quad (9)$$

c) Cl – Ca type III water:

When:

$$\text{Na/Cl} = 0.75 - 0.65 \text{ or likely } 0.60 \quad (10)$$

d) Cl – Ca type IV water:

When:

$$\text{Na/Cl} = 0.65 - 0.50 \quad (11)$$

e) Cl – Ca type V water:

When:

$$\text{Na/Cl} < 0.50 \quad (12)$$

7. Results and discussion:

The aqueous solubility of petroleum hydrocarbons increases with increasing temperature and pressure and decreases with increasing water salinity. Occurrences of petroleum accumulations often correlate

with salinity transition zones, i.e., where the salinity ranges from 50,000 to 100,000 mg/liter (Collins 1975).

Figure (4) illustrates the variation values in Total Dissolved Solids (TDS) concentrations at (mg/l) of oil field waters in Amal Oil – Field, which ranged from 205487 to 234912 mg/l. According to Gorrell's classification, all the studied samples are brine waters. Table (2).

This map showed the changes of TDS in the study area with the contour interval 5000, where TDS values increased in the east and north directions.

Finally, the TDS map illustrates water movement in East and North directions in addition, other words trending from West and South East. Where decreasing at south at (B118 and B86) wells.

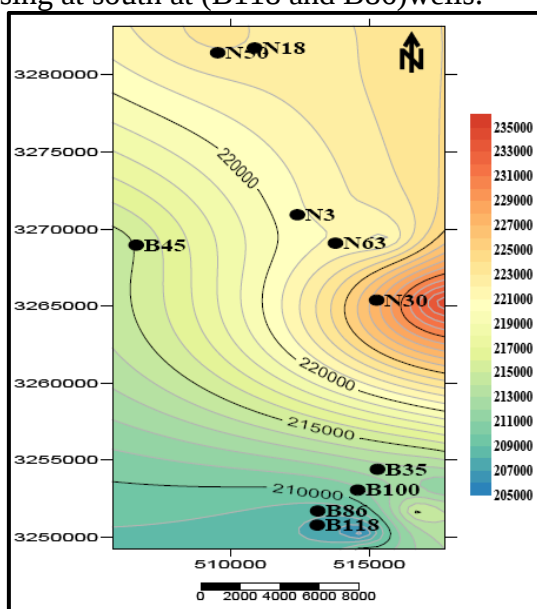


Figure (4) Geochemical map showing distribution of TDS (mg/l) in Amal Oil field.

7.1 Classification of oil field waters:

Water analyses have generally been classified according to a system proposed by Palmer (1911, in Collins, 1975). First, the strong bases (Na^+ , K^+) are combined with the strong acids (Cl^- , SO_4^{2-}) to form primary salinity (S1). Then the strong bases are combined with the weak acid (HCO_3^-) to form primary alkalinity (A1). Next, the weak bases (Ca^{2+} , Mg^{2+}) are combined with the strong acid (Cl^- , SO_4^{2-}) to form secondary salinity (S2). Finally, the weak acids are combined with

the weak bases to form secondary alkalinity (A2). In fact, such combinations of ions do not take place, and it is hard to relate the different water classes to their history and geology. This classification was not used in this study.

The Russian hydrologist Sulin (1946) proposed a classification that is claimed to have genetic significance. This classification of the subsurface hydrochemical system is based upon various combinations of dissolved salts in the waters, which are divided according to Table (1) and equations (4) through (7). There are four major classes of waters; each one refers to different environments from natural water distribution:

1. Continental (terrestrial) conditions that promote the formation of sulfate waters. Such conditions supply soluble sulfate constituents to the water, and the genetic type of such a water is "sulfate – sodium".
2. Continental conditions that encourage the development of sodium bicarbonate waters are characterized by the genetic type known as "bicarbonate – sodium."
3. Marine condition and the formation of a "chloride – magnesium" type of water.
4. The deep subsurface conditions within the Earth's crust lead to the creation of a "chloride – calcium" type of water.

The first two types are indicative of meteoric and/or artesian waters, while the third pertains to marine environments and evaporite sequences. The fourth type is representative of deep stagnant conditions.

Sulin's system (1946) underwent modification by Bojarski (1970 in Collins, 1975). This modification aimed to redefine Sulin's classification of water types and provide a new interpretation of the environments linked to each type, as follows:

1. The "bicarbonate-sodium" water type is found in the upper zone of the sedimentation environment, characterized by hydrodynamic conditions that favor the preservation of petroleum and natural gas deposits.

2. The "sulfate-sodium" water type indicates the reaction of sodium with chloride or sulfate.
3. Waters of the "chloride – magnesium" type are characteristic of the transition zone between hydrodynamic areas, which is becoming more hydrostatic in the deeper part of the basin.
4. Water of the "chloride – calcium" type occurs in deeper zones that are isolated from the influence of infiltration waters and are hydrostatic.

Table (1) Major Classes of Water by Sulin's Classification

Types of water		Ratio of concentrations, meq%		
		Na/Cl	(Na - Cl)/SO ₄	(Cl – Na)/Mg
Meteoric	Sulfate sodium	>1	<1	<0
Mixed	Bicarbonate sodium	>1	>1	<0
Connate	Chloride magnesium	<1	<0	<1
Connate	Chloride calcium	<1	<0	>1

A significant variation in the chemical composition of the chloride-calcium type was noted, and this classification was detailed by Bojariski (1970) in Collins (1975) as follows:

a-Class, "**chloride – calcium I**" with a ratio of Na/Cl > 0.85 refers to an active hydrodynamic zone. It is considered a zone of little prospect for the preservation of the hydrocarbon.

b-Class, "**chloride – calcium II**" with a ratio of Na/Cl = 0.85 – 0.75, refers to the transition zone between hydrodynamic zone and stable hydrostatic zone. It is considered a poor zone for hydrocarbon preservation.

c- Class, "**chloride – calcium III**" with a ratio of Na/Cl = 0.75 – 0.65, represents condition for the favorable preservation of hydrocarbon deposits. It is designated as a fairly hydrocarbon preservation environment.

d-Class, "**chloride – calcium IV**" with ratio of Na/Cl = 0.65 – 0.50, represents a complete isolation environment of hydrocarbon accumulation as

well as by the presence of residual water. This class represents a good zone for hydrocarbon preservation.

e-Class, "**chloride – calcium V**" with ratio of Na/Cl < 0.50, refers to the presence of highly altered residual ancient seawater. This type is one of the most likely indicators for the hydrocarbon accumulation zone.

Table(2) chemical analysis of the oil – field waters from Amal Oil .Field NC-12, sirt basin Libya in mg/l

Wel I No.:	Cations(mg/l)				Anions(mg/l)			TDS (mg/l)	p H	Sp. Gr.
	Na ⁺	k ⁺	Ca ⁺⁺	Mg ⁺ +	Cl ⁻	SO ₄ ⁻⁻	HCO ₃ ⁻			
B- 35	591 01	33 0	196 95	131 2	1301 13	5	195	21075 1	6. 0	1.1 4
B- 45	657 24	95 0	121 44	292 8	1322 40	26	271	21428 3	6. 0	1.1 4
B- 86	599 90	63 0	177 55	583	1262 13	35	281	20548 7	5. 3	1.1 4
B- 100	611 37	12 00	109 22	696 2	1350 76	7	172	21547 6	5. 7	1.1 4
B- 118	607 43	10 0	187 98	960	1297 58	1	239	21059 9	5. 8	1.1 4
N-3	640 97	10 80	185 37	131 2	1364 94	39	239	22179 8	5. 7	1.1 5
N- 18	659 70	12 40	110 2	124 17	1411 03	12 4	166	22212 2	5. 5	1.1 5
N- 30	639 49	12 60	219 84	222 3	1453 57	5	134	23491 2	4. 8	1.1 5
N- 50	593 33	11 20	236 47	131 2	1382 67	20	122	22382 1	5. 8	1.1 5
N- 63	668 88	12 00	125 25	342 6	1364 94	14 0	137	22081 0	5. 4	1.1 5

7.2Classification of study area:

According to sulin, based on the previous formulas from ten samples, all of the waters were Cl – Ca type waters; these results are represented in Table (4). This was done using equations (4) through (7).

In addition, Cl – Ca waters reflect the fact that these waters are associated with oil in a deep subsurface environment. So we can determine water flow directions in oil fields by tracing such types of waters, especially if we have already Cl – Ca waters in any studied field. The analyzed water was classified based on Na/Cl ratios according to the Bojarski (1970) classification modified from the sulin classification. This was done using equations (8) through (12). The results, as shown in Table (5), showed that in the two wells (B – 45, N – 63) the Na/Cl ratios were 0.77 and 0.76, respectively. This means that the water is of the second type of chloride calcium (Cl – Ca II), which indicates that the transitional zone between an active hydrodynamic zone and a more stable hydrostatic zone of the sedimentation basin is considered a poor zone for preserving hydrocarbons.

Regarding the wells (B – 35, B – 86, B – 100, B – 118, N – 3, N – 18, N – 30, and N – 50), the water is identified as type III calcium chloride. This region exhibits favorable conditions for the preservation of hydrocarbon deposits, suggesting that the environment is somewhat conducive to hydrocarbon preservation.

Table(3) chemical analysis of the oil – field waters from Amal Oil Field NC-12, sirt basin Libya in epm.

Well No.:	epm						
	Na ⁺	k ⁺	Ca ⁺⁺	Mg ⁺⁺	Cl ⁻	SO4 ⁻⁻	HCO ³⁻
B-35 1	2251.	7.4	860.6	94.5	3213.7	0.1	2.8
B-45 6	2505.	21.3	531.1	211.	3269.1	0.5	3.9
B-86 0	2289.	14.1	777.2	42.1	3122.8	0.6	4.0
B-100 6	2326.	26.9	476.8	501.	3333.3	0.1	2.5
B-118 7	2315.	2.2	822.1	69.2	3207.7	0.0	3.4

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N-3	2435. 0	24.1	807.9	94.3	3362.4	0.7	3.4
N-18	2504. 0	27.7	48.0	891. 6	3473.0	2.3	2.4
N-30	2412. 5	28.0	951.4	158. 7	3555.9	0.1	1.9
N-50	2252. 0	25.0	1029. 7	94.2	3403.1	0.4	1.7
N-63	2538. 8	26.8	545.4	246. 0	3359.5	2.5	2.0

Table (4) Classification of the oil – field waters from Amal Oil Field NC-12, Sirt Basin Libya, According to Sulin's (1946).

Well No.:	Na⁺ \ Cl⁻ epm	(Na⁺ - Cl⁻) \ SO₄²⁻ epm	(Cl⁻ - Na⁺) \ Mg⁺⁺ epm	Type of water
B-35	0.70	-10559.24	10.18	Cl-Ca
B-45	0.77	-1609.28	3.62	Cl-Ca
B-86	0.73	-1304.45	19.81	Cl-Ca
B-100	0.70	-7895.17	2.01	Cl-Ca
B-118	0.72	-48884.93	12.88	Cl-Ca
N-3	0.72	-1307.78	9.84	Cl-Ca
N-18	0.72	-430.12	1.09	Cl-Ca
N-30	0.68	-12663.88	7.21	Cl-Ca
N-50	0.66	-3167.90	12.22	Cl-Ca
N-63	0.76	-322.66	3.34	Cl-Ca

Libya, According to Bojarski (1970).

Well No.:	Na⁺ \ Cl⁻ epm	Type of water	Class
B-35	0.70	Cl-Ca	III
B-45	0.77	Cl-Ca	II
B-86	0.73	Cl-Ca	III

B-100	0.70	Cl-Ca	III
B-118	0.72	Cl-Ca	III
N-3	0.72	Cl-Ca	III
N-18	0.72	Cl-Ca	III
N-30	0.68	Cl-Ca	III
N-50	0.66	Cl-Ca	III
N-63	0.76	Cl-Ca	II

We observed a notable decrease in sulfate concentrations in the study area, as shown in Table (2). The sulin classification (1946) highlights groups with highly mineralized chloride-calcium or bicarbonate-sodium water types, which contain reduced forms of sulfur. These are significant indirect indicators of oil presence. A low sulfate content is essential to suggest interaction with bituminous components and/or sulfate-reducing bacteria.

8. Graphic representation:

Water analyses are often expressed graphically. The diagram, or pattern, obtained by graphically plotting the results of a water analysis will often highlight important points about the analysis that might be missed by simply reading the report. Pattern comparison is also an easy way to quickly spot differences in two or more waters (Patton, 1986). There are many different water analysis diagrams in use. Two such diagrams are:

8.1 Piper Diagram:

A piper diagram is a graphical representation of the chemistry of a water sample or samples. It shows the relative concentrations of six to seven ions in solutions, in this case, the cations Ca^{2+} , Mg , and $\text{Na}^+ + \text{K}^+$, and the anions Cl^- , SO_4^{2-} , and HCO_3^- . In most natural waters, these ions make up 95 to 100% of the ions in solution.

The Piper diagram includes two trilinear diagrams, one for anions (on the lower right) and one for cations (on the lower left). For each sample, the information from each trilinear diagram is projected up into the central quadrilateral. Therefore, each sample will plot in each frame of the piper, once representing cations, once representing anions, and once representing the combination.

For each constituent, the concentration (in mg/l) is converted to chemical equivalents (meq/l) based on the valence and atomic weight. Then the percentages of each ion relative to the total are calculated and

plotted on the piper diagram. Each trilinear diagram shows the relative percentages of the three ions. Each corner on the triangles represents 100% of the ion shown at that corner. (C. Sadashivaiah et al. 2008). Chemical data of representative samples from the study areas presented by plotting them on a Piper-tri-linear diagram for these wells, as shown in Figure (5), by the geochemical software RockWorks. The piper diagram in the study area indicates most of the samples were classified as $(\text{Na}^+ + \text{K}^+)$ and Ca^{2+} types of cations and Cl^- types of anions. Water in the Amaloil field consists of alkalis that exceed alkaline earths. The dominant cation found in the water is $(\text{Na}^+ + \text{K}^+)$ followed by Ca^{2+} and the dominant anions are Cl^- followed by HCO_3^- .

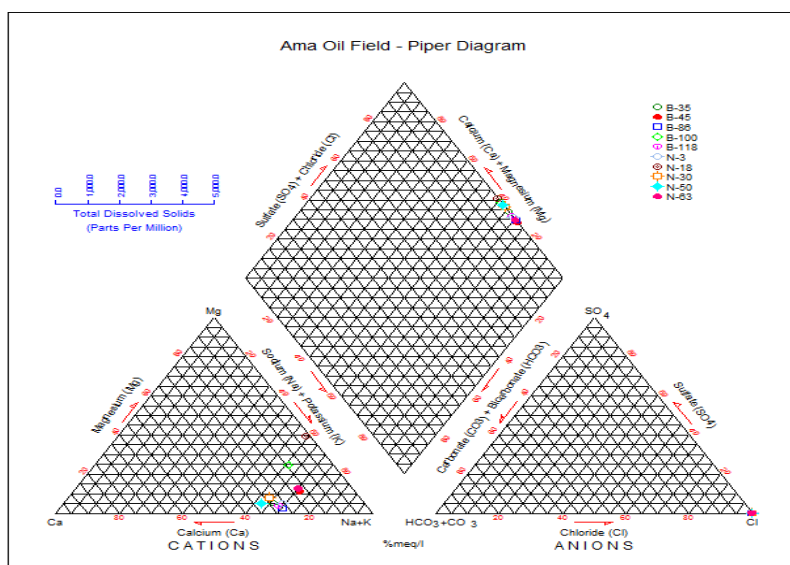


Figure (5) Piper diagram showing proportion of major ions in Amal Oil Field.

8.2 Stiff Diagram:

It is also called a **Butterfly Diagram**. It is a graphical representation of chemical analyses, first developed by H.A. Stiff in 1951. It is widely used by hydrogeologists and geochemists to display the major ion composition of a water sample. A polygonal shape is created from four parallel horizontal axes extending on either side of a vertical zero axes. Cations are plotted in meq/l on the left side of the zero axes, one to each horizontal axis, and anions are plotted on the right side. Stiff patterns

are useful in making a rapid visual comparison between water from different sources, i.e., connate injection waters.

This diagram is the most widely used in the graphic representation of oil-field water analysis. The scale may be either arithmetic or logarithmic. See also Figures 6.a and 6.b (Patton, 1986).

Figures 6.a and 6.b show the stiff diagram for some wells in the Amal oil field; Figure 6a shows the similarity in chemical composition of these waters from the study area. The concentration of chlorine and sodium is high in all wells, followed by magnesium, bicarbonate, and very low amounts of sulfate concentrations; this is because bacteria reduce sulfate to sulfide in a confined environment, reducing the sulfate and bicarbonate contents in the formation waters (Lie et al., 2024).

While we observe that wells B-45 and B-100 have higher concentrations of magnesium and lower concentrations of calcium relative to other wells depicted in the same figure.

As for Figure (6b), we notice that the water in wells N-3, N-30, N-50, and N-63 has almost the same composition, as it contains high concentrations of chlorine and sodium, followed by calcium and then magnesium, which indicate the same origin and source of water. In well N-18, we see a noticeable decrease in calcium concentrations with an increase in magnesium concentrations compared to the rest of the wells in the same figure.

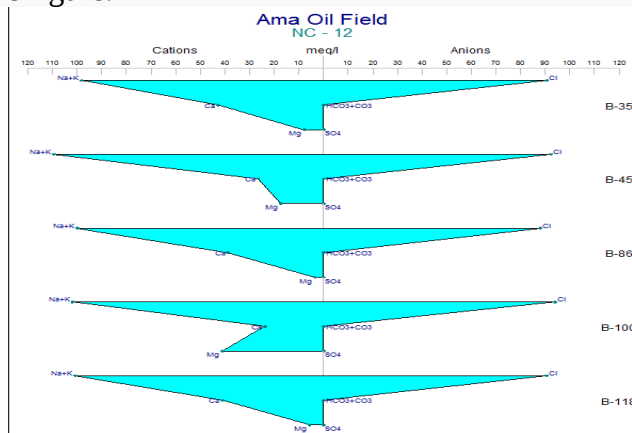


Figure (6.a) Stiff diagram showing the major ion composition of water in the Amal Oil Field.

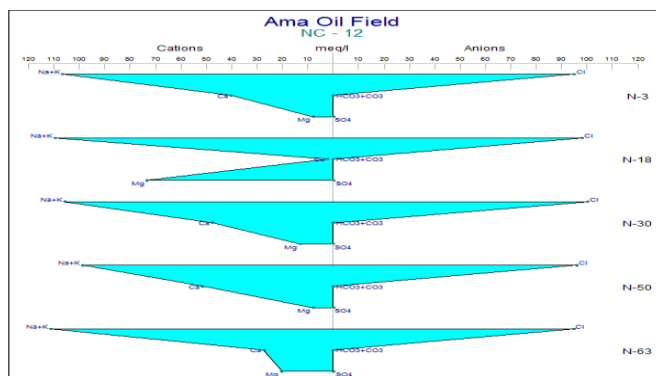


Figure (6.b) Stiff diagram showing the major ion composition of water in the Amal Oil Field

9. Conclusion

This study presents the results of the geochemical classification of some wells in Amal Oil Field in the Sirt Basin, Libya. These wells are: (B – 35, B – 45, B – 86, B – 100, B – 118, N -3, N – 18, N – 30, N -50, and N – 63).

From geochemical calculations and interpretations of the studied oil – field water samples, the following conclusions are illustrated:

Cations in oil – field waters are: Na^+ , K^+ , Ca^{2+} , and Mg^{2+} . Anions are: Cl^- , HCO_3^- , and SO_4^{2-} .

Total Dissolved Solids (TDS) range from 205487 to 234912 mg/l. According to Gorrell's classification, all the studied samples are Brine waters.

According to Sulin's classification, all studied wells are: Cl – Ca type waters (Connate); Cl – Ca waters reflect that these water are associated with oil in a deep subsurface environment and also reflect isolated waters with hydrostatic conditions. These waters were also classified according to the Bojarski, and the results showed that the majority of waters were classified as Cl – Ca type III. This type describes conditions that are conducive to the preservation of hydrocarbon deposits. It is identified as a relatively favorable environment for the preservation of hydrocarbons. Followed by Cl – Ca type II. This area represents the transitional zone between an active hydrodynamic zone and a more stable hydrostatic zone within a sedimentation basin, which is regarded as a less favorable environment for hydrocarbon preservation.

Chemical data of representative samples from the study oil field were presented by plotting them on Piper and Stiff diagrams as follows:

In the piper diagram, most of the samples were classified as ($\text{Na}^+ + \text{K}^+$) and Ca^{2+} types of cations and Cl^- types of anions. Water in the Amal oil field consists of alkalis that exceed alkaline earths, and strong acids exceed weak acids. The dominant cation found in the water is ($\text{Na}^+ + \text{K}^+$) followed by Ca^{2+} and the dominant anions are Cl^- followed by HCO_3^- . Stiff diagrams show increasing concentrations of Na^+ , Ca^{2+} , and Cl^- and decreasing concentrations of Mg^{2+} and SO_4^{2-} . In most wells, which indicate the same origin and source of water, other samples show higher concentrations of Mg^{2+} and lower concentrations of Ca^{2+} . and we note a decrease in SO_4^{2-} concentrations in all wells. These are significant indirect indicators of oil presence. A low sulfate content is essential to suggest interaction with bituminous components and/or sulfate-reducing bacteria. While the increase in magnesium concentration in some wells in the study area is due to the rock description of the formations forming that area, which is represented by the presence of interferences of volcanic rocks with sandstone rocks, as well as perhaps the mineral composition of the reservoir rock, as it was previously indicated that it contains clay and mica minerals.

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